Structural Controls on the Chemistry and Output of the Wells in the Olkaria Geothermal Field, Kenya

(Wamalwa, R.N.)
KenGen
785, Naivasha
rwamalwa@kengen.co.ke

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ABSTRACT
Structures in the Greater Olkaria volcanic complex include; the ring structure, the Ol’Njorowa gorge, the ENE-WSW Olkaria fault and N-S, NNE-SSW, NW-SE and WNW-ESE trending faults. The faults are more prominent in the East, Northeast and West Olkaria fields but are scarce in the Olkaria Domes area, possibly due to the thick pyroclastics cover. The NW-SE and WNW-ESE faults are thought to be the oldest and are associated with the development of the rift. The most prominent of these faults is the Gorge Farm fault, which bounds the geothermal fields in the northeastern part and extends to the Olkaria Domes area. The most recent structures are the N-S and the NNE-SSW faults. The geochemistry and output of the wells cut by these faults have a distinct characteristic that is the N-S, NW-SE and WNW-ESE faults are characterized by wells that have high Cl contents, temperatures and are good producers whereas the NE-SW faults, the Ring Structure and the Ol’Njorowa gorge appear to carry cool dilute waters with less chloride concentration and thus low performing wells. Though the impacts of these faults are apparent, there exists a gap in knowledge on how wide is the impact of these faults on the chemistry and performance of the wells. This paper therefore seeks to bridge this gap by analysis of the chemical trends of both old wells and newly drilled ones to evaluate the impacts of individual faults and then using buffering technique of ArcGis estimate how far and wide the influence of the faults is.

1. INTRODUCTION
Development of geothermal resources in the Olkaria area, a high temperature field, started in the early 1950s. Subsequent years saw numerous expansions with additional power plants being installed in Olkaria (Opondo, 2008). These include a binary plant at Olkaria South West (Olkaria III) in 2000, with a capacity of 110 Megawatts (MW), a condensing plant at Olkaria North East (Olkaria II) in 2003, with a capacity of 105 MWe and another binary plant at Olkaria North West (Oserian) in 2004, with a capacity of 2 MWe. Plans are underway to commission 280 Mw within Olkaria 1 and Olkaria Domes.

Geochemistry in such a setting, given in terms of concentrations of various chemical constituents in both brine and steam discharged from geothermal wells is vital. This is through the understanding that various chemical constituents in both brine and steam discharged from geothermal wells, provide valuable quantitative information that can be used to note the reservoir condition and changes in response to steam exploitation. This is by grouping the elements into two i.e. tracer/conservative elements to trace the origin of the fluids and geo-indicators to reflect the specific environments of equilibrium. The reduction in the concentration of conservative elements will mean dilution occurring in the reservoir leading to cooling hence reduction in the steam production; steam meant to run the turbines for electricity generation. The geo-indicators which are chemically reactive reflect specific environments of equilibrium. For example Sodium, Silica and Potassium to be used as solute Geothermometers. Such studies apart from being used directly in reservoir modeling and simulations, they also assist with calibration of the other reservoir models as trends in physical and chemical parameters change.

2. Structural Geology of Olkaria and profiles selection

2.1 Structural Geology of Olkaria and Chemical Data selection
Geological structures within the Greater Olkaria volcanic complex include; the ring structure, rift fault systems, the Ol-Njorowa gorge, and dykes swarm (Fig. 1). The faults are trending ENE-WSW, N-S, NNE-SSW, NW-SE and WNW-ESE. These faults are more prominent in the East, Northeast and Olkaria West fields but are elusive in the Olkaria Domes area, possibly due to the thick cover of the surface by pyroclasts (Lagat, 2004).

The NW-SE and WNW-ESE faults are thought to be the oldest fault system and they link the parallel rift basins to the main extensional zone (Wheeler and Karson, 1994). Gorge Farm fault, bounding the geothermal fields in the north eastern part and extending to Olkaria Domes area is the most prominent of these faults.

The most recent structures are the N-S (Ololbutot eruptive fissure) and the NNE-SSW faults (Mungaia, 1999). Dyke swarms exposed in the Ol-Njorowa gorge with strike direction in the NNE further attest to the recent reactivation of faults with this trend. Development of the Ol-Njorowa gorge was initiated by faulting along the trend of the gorge but the feature, as it is seen today, was mainly formed due to a catastrophic outflow of Lake Naivasha during its high stands (Clarke et al.,...
Volcanic plugs (necks) and felsic dykes occurring along the gorge further attest to the fault control in the development of this feature. Hydroclastic craters located on the northern edge of the Olkaria Domes area mark magmatic explosions, which occurred in submerged country (Mungania, 1999). These craters form a row along, which the extrapolated caldera rim trace passes.

Figure 1: Structural map of the Greater Olkaria Geothermal Area (Omenda, 1998).
2.1.2 Profiles Selection

Study considered 6 profiles along the structures cutting through the field as shown in Fig. 2 below.

![OLKARIA GEOTHERMAL FIELD SECTORS AND THE SELECTED PROFILES](image_url)

**Figure 2:** Olkaria Geothermal Sectors and Selected Profiles.

The profiles allowed a distance of 1 km wide to cater for the buffering process.

3. Chemical Data collection

The wells located in the production are usually monitored through geochemical sampling programs. The east production field, for example, is monitored in a biannual program in which gas samples and water samples are collected. The North east production field is monitored on a quarterly basis under the program of clean development mechanism (CDM) in which gas and water samples is collected for analysis. Newly drilled wells in the various sectors of development or further development have been discharge tested fluid chemistry of the areas made available.

3.1 Chemical Sampling

The sampling and analysis methods are based on those of Arnasson et al. (2007) where steam samples were collected with a stainless steel webre separator connected between the wellhead and an atmospheric silencer on the discharging wells. For the earliest wells in the Domes, only the steam in the two phase flow was sampled. The water samples were taken from the weir box down the stream from an atmospheric silencer. However this changed later. The later sampling started collecting both the steam and water from the two phase flow from the line between the wellhead and the atmospheric silencer using a webre separator. Another water sample was also obtained from the weirbox.

3.2 Chemical Analysis

The analysis is done in house in the KenGen geochemistry laboratory according to standard methods for geothermal fluid analysis as described by various authors (Arnórsson et al., 2000 and Arnórsson et al., 2007). It is desirable that the charge balance error of geothermal waters be within ±5%. Therefore data that never met this criterion was left out.

4 Geochemical Variations

The study considered one conservative element (Cl), one geoindicator (TNaK) and nitrogen gas to help in delineating the source of fluids, reservoir temperatures and signs of incursion of atmospherically modified fluids respectively. The results are in form of graphs of concentration variation with time.
Figure 3: Chloride Variation With time per profile.
Figure 4: Temperature Variation with time per profile.
PROFILE 1 CHEMICAL VARIATION

Dates of Discharge

Years of Production

Concentration (mm/100mole)

PROFILE 2 CHEMICAL VARIATION

Days of Discharge

Years of Production

Concentration (ppm)

PROFILE 3 CHEMICAL VARIATION

Days of Discharge

Years of Production

Concentration (ppm)

PROFILE 4 CHEMICAL CHANGES

Years of Production

Concentration (ppm)
5 Discussions
This is in form of distribution maps plotted per profile (Fig. 5, 6 and 7). The Chloride and temperature distribution maps indicate the NW-SE trending faults bring in Cl rich and high temperature waters; these increase towards the central field a factor that has been attributed to boiling due pressure drop arising from production. Profile 5 and 6 give a clear distinguishing factor in that wells around the Ol Njorowa gorge, which has majorly NE-SW trending faults have low Cl content whereas those cut by NW-SE faults like OW-44 and OW-915 has high contents.
Figure 6: CI Distribution Map per profile

Figure 7: Temperature Distribution per profile

This study also considered N$_2$ as it is the best indicator of incursion of atmospherically modified reservoir fluid. The distribution map (Fig. 8) show low modifications associated with the NW-SE trending faults. As in case of profile 1, there is a sharp increase evident towards the SE corner; areas bordering the Ol Njorowa gorge. This therefore indicate that this feature act to recharge the reservoir.
Figure 8: Nitrogen Distribution per profile
6. Enthalpy Variation in the Producing wells

The enthalpy variation has been used to summarize the wells output with time (Fig. 9). The trending show that the NW-SE (Olkaria fracture), bringing cooler fluids has impacted on OW-725/ OW-705 (see profile 3&4 Distribution map). These wells have the least combined enthalpies in the Olkaria North East field. The effect of the Ol Njorowa Gorge is also evident with the declining enthalpy and consequently output of OW-16, OW-19, OW-22 and OW-23 (see Profile 5 & 6 Distribution map).

![Figure 9: Enthalpy Variation with time](image)

7 Conclusions

After the evaluation of the chemical data, it was evident that Olkaria field has three categories of faults that control fluid flow:

I. The NW-SE trending faults that bring in high temperature and Cl rich waters.

II. The NE-SW trending Olkaria fracture tend to carry cool temperature waters that has led to decline in enthalpies of the wells it cuts through

III. The faults within the Ol Njorowa gorge act to carry cool, less mineralized water. Though initially, these effects were thought to be in shallow depths, an indication in OW-901 which is a deeper at 2200 m compared to 1600 m of OW-23 well proves otherwise. This is however to be proved later as much deeper wells have been sited.
REFERENCES


